

REVISED PALEOGEOGRAPHY OF THE DINOSAUR BEARING MAASTRICHTIAN LAMETA FORMATION, CENTRAL AND WESTERN INDIA: IN THE PERSPECTIVE OF NEWLY IDENTIFIED SALBARDI-BELKHER INLAND BASIN

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Abstract

The Lameta sediments covering the geographical area of about 5,000 km2 in central and western India are well documented for their sedimentological and paleobiologic aspects in last four decades. Occurrences of dinosaurian remains including bones, coprolites, eggs, eggshells and nests in mainly arenaceous lithounits of the successions exposed at various localities have been widely used as a significant tool for precise reconstructions of paleogeographic and paleoenvironmental set-up during the Maastrichtian period. Based on sedimentological and paleobiologic aspects, earlier it was considered that the Lameta sedimentation took place in only five inland basins viz., i) Nand-Dongargaon, ii) Jabalpur, iii) Sagar (Saugor), iv) Ambikapur-Amarkantak and v) Balasinor-Jhabua. Recently, a sixth inland basin has been added by Mankar and Srivastava, Salbardi-Belkher, which has been highlighted because of its bearing on the overall picture of Lameta sediments deposition, dinosaur inhabitation and its paleogeographic limit. The categorizations of these inland

1. Introduction

The Lameta sediments were widely reported from central and western parts of the India exposed at district Jabalpur, Madhya Pradesh, Nagpur and Chandrapur districts of Maharashtra (Tandon et al., 1990, 1995; Mohabey et al., 1993; Mohabey, 1996 a; Khosla et al., 2005; Mohabey and Samant, 2005, 2009, 2013; Fernández and Khosla, 2015) and Kheda districts of Gujarat (Mohabey, 1984; Srivastava et al., 1986; Sahni, 1993; Khosla and Sahni, 1995, 2003; Mohabey and Samant, 2013; Khosla and Verma, 2015). Scattered occurrences were also reported from Sagar and Amrakanthak districts of Madhya Pradesh (Mohabey, 1996 a) and adjoining areas of Amravati districts of Maharashtra Citation:

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basins had been done on the basis of detailed study including correlations of litho- and biofacies in respective areas. The presently explored Salbardi-Belkher area, owing several localities of Lameta sediments having preservations of skeletal remains, nest and eggs of dinosaurs, is altogether a separate geographical set-up, besides the earlier reported five inland basins. Taking into consideration of the lithofacies architecture of various successions of the new basin, it has been interpreted that the deposition of the Lameta sediments during the Maastrichtian period was contemporaneously taking place in this additional inland basin, situated in the west of Nand-Dongargaon inland basin. This addition of new inland basin of fluvial regime including evidences of dinosaur remains depicts a revised paleogeographic set-up of Lameta sedimentation in central and western India.

Keywords: Paleogeography. Maastrichtian. Dinosaur. Inland Basin. Fluvial-Lacustrine. Paleoenvironment.

and Betul, Madhya Pradesh (Srivastava and Mankar, 2012 a, b, 2013, 2015 a, b; Mankar and Srivastava, 2015) (Fig. 1). These successions were mostly considered to be the deposits of fluvial-lacustrine set-up; however, the column exposed at the type area of Jabalpur bears a debate about its environment of deposition i.e., coastal complex shallow marine vs. fluvial-lacustrine. The former option has been proposed on the basis of the occurrence of algal structures and glauconite minerals in green colored sandstone bed of the succession (Chanda, 1963 a, b, 1965, 1967; Singh, 1981; Singh and Srivastava, 1981; Chanda and Bhattacharya, 1966); lithological architecture including horizons with extensive development of crab burrows i.e., *Thalassinoides* (Kumar and Tandon, 1977, 1978, 1979).





Fig. 1. Map showing: A) distribution of Lameta sediments in various inland basins viz., i) Nand-Dongargaon, ii) Jabalpur, iii) Sagar, iv) Ambikapur-Amarkantak, v) Balasinor-Jhabua and vi) Salbardi-Belkher and major sea way along the Narmda-Tapi rift (Mohabey, 1996 a; Bajpai, 2009; Keller et al., 2009; Mankar and Srivastava, 2015), B) map showing: study area in a) regional set-up, b) local set-up and c) locations of Bairam, Belkher, Pandhari and Salbardi localities of Lameta (GSI 2001, 2002) (Srivastava and Kandwal, 2016).



Recently, Shukla and Srivastava (2008) and Saha et al. (2010) suggested coastal complex setting on the basis of records of lizard eggs and trace fossils of marine affinity. The other group of researchers favoring fluvial-lacustrine environments of deposition for most of the Lameta sediments in India including the succession of type area, advocates on the basis of fossils biota and lithological characteristics as the main tools i.e., non-marine flora and fauna including abundant skeletal remains, eggs, eggs-nests and coprolites of dinosaurs (Mately, 1921; Pascoe, 1964; Mohabey, 1996 a; Srivastava and Mankar, 2013, 2015 b); characteristic lithological association of alluvial plain environment under hot and semi-arid conditions, Jabalpur area (Brookfield and Sahni, 1987); fluvial and pedogenically modified semi-arid, pre-palustrine flat system, Jabalpur area (Tandon et al., 1990); carbonate-chert fluvio-limnic association in Gujarat State (Joshi and Ganapathi, 1990; Mohabey, 1991); facies association including sedimentary petrological characteristic depicting lacustrine set-up, Nand-Dongargaon area, Nagpur (Singh and Tandon, 2004; Mohabey and Samant, 2005); occurrence of pedogenic calcretes formed in sub-aerially exposed area of low gradient alkaline flat alluvial setting, Jabalpur (Tandon et al., 1995, 1998; Tandon and Andrews, 2001); lithological set-up, facies architecture and granulometric analysis of the succession exposed in Salbardi-Belkher inland basin (Srivastava and Mankar, 2010, 2012 a, b, 2015 a; Mankar and Srivastava, 2015; Srivastava and Kandwal, 2016). All these co-eval exposures of different localities exhibit a remarkable similarity in their lithological set-up as well as the faunal contents. Abundant dinosaurian remains e.g. skeletal fragments, eggs, eggs-nests and coprolites from the Lameta successions of various areas belonging to sauropod and theropod families indicate semi-arid climate and fluviallacustrine environments of deposition i.e., Jabalpur inland basin (Tandon et al., 1995; Mohabey, 2001; D'Emic et al., 2009; Carrano et al., 2010), Nand-Dongargaon (Mohabey and Udhoji, 1993), Balasinor-Jhabua (Mohabey, 1983, 1987; Jain and Bandopadhyay, 1997), Salbardi-Belkher (Srivastava and Mankar, 2013, 2015b).

Based on the lithological characteristics and faunal contents, Mohabey (1996 a) identified five separate inland basins of fluvial-lacustrine nature, in which, the Lameta sediments were deposited simultaneously viz., i) Nand-Dongargaon, ii) Jabalpur, iii) Sagar, iv) Ambikapur-Amarkantak, and v) Balasinor-Jhabua. The architectures of these inland basins show remarkable similarity; however, differs in lateral correlatability as having basinal control instead of time (Mohabey, 1996 a; Shukla and Srivastava, 2008). Recently, a new inland basin of Lameta sedimentation has been proposed on the basis of limited study viz., Salbardi-Belkher, making a total of six inland basins in central and western India (Mankar and Srivastava, 2015). This basin was established on the basis of new record of 7-8 nearby isolated Lameta exposures, lying in the vicinity of

10-40 km aerial distances, of which, four successions of comparatively larger dimensions are exposed at the localities of Bairam (21°22'25" N; 77°37'23" E), Belkher (21°21'48" N; 77°31'23" E), Pandhari (21°22'02" N; 77°32'54" E) and Salbardi (21°25'15" N; 78°00'00" E) areas, in addition to 3-4 small exposures at nearby places. The successions at two localities viz., Salbardi and Pandhari, have preservations of dinosaur bones and eggs. Recently, certain calcareous algae have also been reported from discontinued thin bed of calcmarl, preserved in the lower part of the succession at Pandhari area (Srivastava et al., 2018). The identification of the Salbardi-Belkher basin is based on the establishments of lithological architectures of various successions exposed at above mentioned localities, reconstruction of depositional environments and the presence of dinosaurian remains.

In the present study, an attempt has been made to discuss the Salbardi-Belkher inland basin in detail and to compare its depositional environment set-up and dinosaurian remains with those of the other five basins aiming to establish a regional picture of paleogeographic set-up in central and western parts of the India during Maastrichtian times.

2. Geology, age and stratigraphy of the area

The Lameta successions confined to Salbardi-Belkher inland basin, disconformably overlying along with the Upper Gondwana succession (\approx the Jabalpur Formation), are exposed in the basaltic country of Deccan Trap as an inlier due to tectonic activity of the Satpura and other allied faults (Fig. 1, Table 1). In the regional set-up, quartz-feldspathic gneiss of the Archaean age forms the basement. The Upper Gondwana succession unconformably overlies the Archaean and is mainly represented by areno-argillaceous sediments having preservation of medium to large scale cross and parallel beddings. This sedimentary succession 15 disconformably overlain by the Lameta sediments which are areno-argillaceous and calcareous in nature. Deccan trap forms the capping and represented by melanocratic, hard, massive and vesicular to amygdaloidal basalts. The lithological similarities of the Gondwana sediments of Salbardi area with those of the exposures at Bairam and Belkher localities suggest the same age hence, considered being a co-eval lithounit (Srivastava and Mankar, 2008). The age of the Upper Gondwana successions exposed at Bairam and Belkher area is assigned as Upper Jurassic to Early Cretaceous on the basis of gymnosperm and pteridophytic remains (Srivastava et al., 1995, 1999).

3. Lameta successions of the area

Despite 3-4 small, scattered patches, the basin shows comparatively good development of the Lameta successions at four places viz., Bairam, Belkher, Pandhari and Salbardi. The successions exposed at all these localities are broadly similar in nature; however, lateral facies variation, pinching and swelling of various lithounits have been noticed. At



REVIEW PAPER

Bairam locality, the column is ca. 39 m thick and represented mainly by clay-marl, arenaceous and calcareous sediments of almost equal thickness (Fig. 2). Here, the lower 12 m of the column consists of brownish-yellowish-greenish clays which are occasionally interbedded with thin beds of siliceous limestone and medium to fine grained sandstones. The middle 13 m of the column is mostly arenaceous, of which, the lower 2 m succession is gravish brown, medium to coarse grained sandstone, overlain by 9 m of clay-marl, which contains thin discontinuous beds of siliceous limestone. The overlying 2 m lithounit is hard and compact, medium to coarse grained bioturbated sandstone. This column shows preservation of abundant arenaceous concretions of 2 to 14 cm diameters. The upper 14 m of the column is calcareous, which is easily divisible into nodular limestone and chertified limestone. The nodular limestone contains clasts of chert and Jasper, whereas, the chertified limestone is indurated and shows poor tendency of flat bedding. Chertification is mostly in the form of horizontal discontinued beddings.

The succession outcropping at Belkher area is comparatively well developed and attains the height of about 47 m (Fig. 2). The arenaceous sediments are comparatively more developed and form the lower 21 m of the column with occasional occurrences of clay bands. The sandstones are mostly yellowish orange to grayish brown in color, medium to fine grained and show parallel beddings and cross beddings. Preservations of *Thalassinoides* and other feeding burrows have also been reported from grayish brown sandstone lithounit (Srivastava and Mankar, 2012 a). The middle 12 m of the column is calc-marl and contains abundant greenish black, subangular to subrounded, hard and compact concretions. The upper calcareous column, consisting of nodular and chertified limestones, having a thickness of about 14 m, is similar to that of the Bairam area.

Tab. 1. Regional stratigraphic set-up of the area.

Age	Stratigraphic units	Rock types				
Quaternary		Alluvium and Soil				
	Unconformity					
L. Cretaceous to Eocene	Deccan Trap	Grayish black, hard and compact, non porphyritic and porphyritic basalts				
Unconformity						
L. Cretaceous	Lameta	Sandstone, claystone, mudstone and limestone				
Disconformity						
E. Cretaceous	Upper Gondwana (≈Jabalpur Formation)	Sandstone, siltstone, conglomerate, claystone and mudstone Upper Gondwana				
Unconformity						
Archaean		Quartz-feldspathic gneiss with dolerite intrusions.				

The succession at Pandhari area (Fig. 2), with a thickness of ≈ 37 m can be considered as an eastward extension of the Belkher succession, lying at an aerial distance of about 1.5 km; however, differs locally in lithological set-up. Here, argillaceous lithofacies are well developed in the lower 9 m of the column, represented by greenish gray to yellowish brown clays with irregular sandy concretions and discontinued beds of light to dark gray micritic limestone. The overlying 19 m of the column is dominantly arenaceous, having shades of pale brown to pale reddish, medium to fine grained, friable to hard sandstones. The upper 5 m part of this arenaceous column shows preservation of abundant irregular concretions of 10 to 13 cm diameter, having similar lithology as of the host rock. The top 9 m of the column is calcareous in nature, represented by brecciated, nodular and chertified limestones.

Salbardi area exhibits 35 m thick sediment column, of which, the lower 9 m is dominantly arenaceous and represented by dark brown, thinly bedded, coarse to medium grained sandstone with parallel and cross beddings. Intercalations of clavey beds can also be noticed. The overlying 10 m column exhibits frequent lithological variation i.e., the lower 2 m succession is calcareous and shows pedogenic activity represented by cylindrical calcretes. It is overlain by one-meter thick, light green, parallel bedded calcareous sandstone succeeded upward by 2 m thick light brown micritic calcrete which is overlain by 5 m thick reddish-brown clayey beds. The upper 16 m calcareous unit is represented by nodular limestone and chertified limestone or, occasionally by their alternations. The top 4 m of this calcareous lithounit is represented by intraformational brecciated limestone which is micritic and contains large clasts of nodular and chertified limestones.





Fig. 2. Comparative view of the lithologs showing vertical set-up of Lameta successions at various inland basins in central and western India (Mankar and Srivastava, 2015).

4. Dinosaur remains including bones, eggs and coprolites

Dinosaur remains were collected from the two areas viz., Salbardi and Pandhari. The light green colored sandstone of Salbardi area has abundant fragmentary bones (Fig. 3A-D). A fragment of right ulna belonging to *Titanosaurus colberti* has been reported by Srivastava and Mankar (2013) (Fig. 3E). In addition, nests and eggs belonging to Megaloolithus oogenus of Megaloolithidae oofamily confirm the area as a suitable site for nesting (Srivastava and Mankar, 2015 b) (Fig. 3F). Pandhari area includes fragmentary bones and rare occurrence of coprolites.

5. Facies analysis and depositional environments

Srivastava and Mankar (2010, 2012 b, 2013) have already documented detailed petrography, lithofacies analysis and reconstruction of depositional environments for the successions of Bairam, Belkher and Salbardi areas. In the present paper, a generalized idea about the lithofacies architecture of the succession is being provided. Basically, three lithofacies associations have been identified which are as follows:

5.1. Arenaceous lithofacies association

- 5.1.1. Massive sandstone lithofacies
- 5.1.2. Green sandstone lithofacies
- 5.1.3. Thinly bedded, yellowish orange and grayish brown sandstone lithofacies
- 5.1.4. Coarse grained sandstone lithofacies
- 5.1.5. Dark brown bioturbated sandstone lithofacies
- 5.2. Argillaceous lithofacies association
 - 5.2.1. Yellowish-brownish-greenish clay-siltstone lithofacies

5.2.2. Light gray silty clay with concretions lithofacies

- 5.3. Calcareous lithofacies association
 - 5.3.1. Calcrete lithofacies
 - 5.3.2. Nodular limestone lithofacies
 - 5.3.3. Chertified limestone lithofacies
 - 5.3.4. Intraformational breccia lithofacies

Table 2 shows the lithofacies architecture of the succession, various lithofacies identified, their characteristic features and depositional environments. Based on the lithological architecture of various successions exposed at three different localities, the depositional environment setup of the basin has been interpreted in detail and a model showing successive developments in fluvial set-up and deposition of corresponding rock types have also been suggested (Srivastava and Mankar, 2015 a) (Fig. 4). The depositional model depicts an overall fluvial-lacustrine set-up, in which, temporal variations in depositional setting and climatic conditions have also been observed. The lower part of Lameta succession at Salbardi area, overlying disconformably the Gondwana succession, reflects low **JSE** REVIEW PAPER

energy condition of deposition. The presence of rounded to subrounded pebbles in both stratigraphic units depicts their common source. Hydrodynamic fluctuations that cause changes in the nature and granulometry of the deposit are well represented by the deposition of fine-grained argillaceous sediments and sandstone with cross and parallel beddings. Seasonal variability of hydrodynamics of the river system is marked by the formation of pedogenic calcretes during arid climate, detached lakes and water bodies during low water condition followed by over flooding and highwater condition of the river channel. A calm and quite water environment is interpreted for a short time period represented by bioturbated sandstone indicating high growth of benthonic fauna and complete churning of the sediments by animal community (Srivastava and Mankar, 2015 a).

The upper part of the succession shows a major change allowing the deposition of predominantly carbonate sediments in an alkaline environment. The huge deposition of nodular and chertified limestones indicates a sheet flood environment of deposition. This phase shows termination of Lameta sedimentation. In the Salbardi area, an additional lithounit of intraformational breccia has been recorded which contains clasts of preexisting rocks in micritic matrix showing gravity flow mode of sediment transport.

6. Lameta Formation of other inland basins

6.1 Jabalpur inland basin

The basin incorporating the Lameta exposures around Jabalpur area includes the localities of Lameta Ghat, Chui Hill and Bara Simla Hill. The first one represents the type area of the Lameta Formation and is well known for various sedimentological and paleobiological studies; however, the other two localities are also well documented. In the type area, five major lithounits have been reported i.e., i) Green sandstone, ii) Lower limestone, iii) Mottled nodular marl, iv) Upper limestone and v) Upper sandstone (Fig. 2, Table 3) (Tandon et al., 1995). The succession exposed at Lameta Ghat is 18-21 m thick and represented by Lower limestone, Mottled nodular marl and Upper limestone (Singh, 1981; Saha et al., 2010). The adjacent area of Bara Simla Hill exhibits all the five major lithounits constituting 34 m thick column (Singh, 1981; Singh et al., 1983), whereas, at Chui Hill, it was earlier reported to be 39 m thick by Singh (1981) which was further modified by Saha et al. (2010) as 45 m thick and absence of the Upper sandstone lithounit. Singh (1981) interpreted the successions as the deposits of estuarine channel to shallow marine coastal complex which is mainly based on the occurrence of Green sandstone and trace fossils. The shallow marine environment has been suggested by various researchers taking into consideration the petrology, lithological architecture, algal structures,



glauconitic beds and extensive crab burrows including *Thalassinoides* (Chanda, 1963 a, b, 1965, 1967; Kumar and Tandon, 1977, 1978, 1979; Singh 1981; Singh and Srivastava, 1981).

Shukla and Srivastava (2008) and Saha et al. (2010) reported trace fossils and lizard eggs from the Lameta Formation of Jabalpur area. The trace fossils represented by *Arenicolites, Zoophycos, Rhizocorallium* and *Ophiomorpha* depicting typical marine environment are recorded from the Lower limestone, Mottled nodular limestone and Upper

limestone. On the basis of ichnofacies, lithofacies architecture and nesting habit of lizards, they interpreted the succession to be a product of intertidal to supratidal channels, marsh, estuary and lagoon subenvironments.

Recently, Srivastava et al. (2015) reexamine the eggshell fragments previously assigned to lizard eggs (Shukla and Srivastava, 2008) and reassigned them to crocodiles and interpreted the paleoenvironment as near shore, lagoonal and supratidal settings of deposition.



Fig. 3. Photographs showing: A-D) fragments of dinosaur bones, E) reconstruction of right ulna of *Titanosaurus colbarti* (Srivastava and Mankar, 2013) and F) spherical eggs belonging to Magaloolithus oogenus (Srivastava and Mankar, 2015 b) (all from Salbardi area).

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REVIEW PAPER

Tab. 2. Comprehensive view of various lithofacies identified, their characteristic features and respective depositional environments for the Lameta successions of Salbardi-Belkher inland basin (revised after Srivastava and Manakar, 2013).

Facies association	Lihtofacies	Characteristics features	Depositional environment
Arenaceous lithofacies association	Massive sandstone lithofacies	Medium grained, hard and compact lithounit with ferruginous cementing material. Parallel beddings and cross beddings are difficult to record. Lithounit contains pebbles similar as of underlying Gondwana succession, however, less in number. It indicates that the source of sediments for both Gondwana and Lameta formations were the same.	Fluvial (channel floor)
	Green sandstone lithofacies	Light to medium green, hard and compact, medium grained lithounit having calcareous cementing material. Large scales cross beddings and parallel beddings are well preserved which indicate high energy condition of deposition. Alternations of thin and thick beds of sandstone having large scale and small-scale cross beddings showing fluctuations in the energy condition of the depositional environment. This is also evident by the truncation of structures due to erosional activity.	Fluvial (point bar)
	Thinly bedded, yellow orange and grayish brown sandstone lithofacies	Alternations of yellowish orange and grayish brown sandstones consisting of medium to coarse grained sediments. Grayish brown sandstones have ferruginous cementing material whereas, siliceous to ferruginous in yellowish orange sandstones. Presence of large-scale cross beddings and rounded to subrounded pebbles indicate medium to high energy condition of deposition.	Fluvial (point bar)
	Coarse grained sandstone lithofacies	Poorly sorted, coarse to medium grained sandstone with abundant granule-size clasts of subangular to angular quartz and feldspar. It indicates a local variation in the energy condition and the influx of sediments from the nearby area.	Fluvial
	Dark brown, bioturbated sandstone lithofacies	Hard and compact, dark brown, medium grained sandstone showing nodular tendency. The unit also shows high bioturbation and lack of mechanically formed sedimentary structures. Extensive development of crustacean community animals indicates well oxidized environment.	Fluvial
Argillaceous lithofacies association	Yellowish-brownish- greenish clay-siltstone lithofacies	Dominance of clayey argillaceous sediments with thin and poorly laminated silty beds. Often, the clay with increasing content of silt size sediments grades to siltstone. Occurrence of abundant calc–marl concretions. The lithounit shows a quiet shallow aquatic environment allowing the suspension of fine sediments.	Lacustrine
	Light gray silty clay with concretions lithofacies	Contains abundant calc-marl concretions having scattered occurrence in the lithounit. Indicates a restricted, comparatively deeper water environment with low energy condition. Presence of angular sand size grains revels the influx of sediments from nearby source.	Lacustrine



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Tab. 2. (cont.) – Comprehensive view of various lithofacies identified, their characteristic features and respective depositional environments for the Lameta successions of Salbardi-Belkher inland basin (revised after Srivastava and Manakar, 2013).

Facies association	Lihtofacies	Characteristics features	Depositional environment
Calcareous lithofacies association	Calcrete lithofacies	Hard and compact, grayish white to light gray calcretised bed with nodular and root calcrete. Both host rock and calcretes include clasts of chert, basalt and green sandstone. The lithofacies is interpreted to be a product of subaerially exposed, low gradient, alkaline flat environment.	Pedogenic
	Nodular limestone lithofacies	Hard and compact, micritic limestone having subangular to subrounded, medium to coarse, detrital grains and, angular to subangular clasts of chert, jasper, feldspar, quartz etc. up to the dimension of 3 cm. Beds show the nodular tendency. Lack of sedimentary structures. On the basis of micrite as the dominant constituent and lack of internal bedding structures, it is interpreted to be a deposit of alkaline flat environment having low energy condition.	Flood plain
	Chertified limestone lithofacies	Hard and compact chertified miciritic limestone. Sedimentary structures are mostly defused, but beds and lenses of chert are present however, the chertification increases in upward direction. It is interpreted to be the continuation of depositional environment as of the previous lithofacies; however, with reduced alkalinity which may be due to the influx of the siliceous dust produced by the volcanic eruption.	Flood plain
	Intraformational breccia lithofacies	Hard and compact, brecciated lithounit having abundant, large sized angular to subangular clasts of nodular limestone, chertified limestone and basalt. Irregular occurrences of angular clasts of preceding lithologies in the micritic groundmass show its deposition by sediment gravity flow. It might be a localized phenomenon.	Gravity flow

The environment of deposition for the successions of this basin was also interpreted to be fluvial-lacustrine by many other researchers. This fluvial-lacustrine environment was suggested mainly on the basis of lithological characteristics (Tandon et al., 1995), occurrence of various types of calcretes (Brookfield and Sahni, 1987; Tandon et al., 1990, 1995, 1998; Tandon and Andrews, 2001) and dinosaurian remains (Matley, 1921; Pascoe, 1964; D'Emic et al., 2009).

6.1.1. Dinosaur remains

Rich assemblage of dinosaur skeletal remains was reported from the Jabalpur area i.e., caudal vertebrae, tooth and skull of sauropods, theropods and ornithopods (Table 4). Sauropods include *Titanosaurus indicus* and *Antarctosaurus* septentrionalis; theropods are represented by *Indosuchus* (Indosaurus) raptorius, Indosuchus matleyi, Lametasaurus indicus, Composuchus solus, Laevisuchus indicus, Jubbalppuria tenuis, Dryptosauroides (?) grandis, Ornithomimoides mobilis, Ornithomimoides barasimlensis; and ornithopoda includes Brachypodasaurus gravis (Hislop, 1859; Lydekker, 1877, 1879, 1890; Chatterjee, 1978; Mathur and Pant, 1986; Mohabey, 1987; Vianey-Liaud et al., 1987; Huene and Matley, 1993; Jain and Bandhopadhya, 1997; Loyal et al. 1999) (Table 4). Recently, Wilson et al. (2003) and Wilson and Upchurch (2003) reexamined various species reported from Lametas of India on the basis of valid taxonomic characteristics and assigned valid species names for sauropods and theropods. Accordingly, the valid species of sauropods includes *Titanosaurus colbarti* (now referred as *Isisaurus colberti*) and *Jainosaurus septentrionalis*; whereas, theropods are limited to only *Inosuchus raptorius, Indosuchus matleyi* and *Laevishuchus indicus*.

6.1.2. Nests and Eggs

Dinosaur nests, eggs and eggshells are well reported from this inland basin. The eggs occur in association with the nest and are well preserved in the form of complete, incomplete



Tab. 3. Lithological set-ups and depositional environments of the Lameta formations exposed in various inland basins of central and western India (Mankar and Srivastava, 2015).

		Jabalpur Basin	Nand-Dongargaon Basin	Balasinor-Jhabua Basin	Salbardi-Belkher Basin	
Lithology 1) U 2) U 3) M 4) L 5) G (Sin		 Upper sandstone Upper limestone Mottled nodular marl Lower limestone Green sandstone (Singh, 1981; Tandon et al., 1990,1995) 	 Red green silty clays associated with sandstone Channel related sandstone with calcretised in upper part Yellow laminated clays and interbedded with limestone, marlite and septarian concretion bands Calcrete Gray nodular marls (Mohabey,1996 a, b; Mohabey and Samant, 2005) 	 Arenaceous limestone Calcareous sandstone Grits Conglomerate (Sarkar et al., 1991; Singh and Tandon, 2004) 	 Intraformational brecciated limestone Chertified limestone Nodular limestone Clay-marl with concretions Yellowish-greenish-reddish clay- marl Bioturbated sandstone Grayish yellowish-brownish sandstone Green sandstone Srivastava and Mankar, 2010, 2012b; 2013a) 	
Depositional environment	Marine	1) Intertidal1) Intertidal-supratidal2) Intertidal2) Intertidal channels on the3) Tidal flat2) Intertidal channels on the4) Tidal flat3) Marsh5) Estuarine4) Lagoon(Singh, 1981)5) Estuary(Saha et al., 2010; Srivastava et al., 2015)				
	Non- marine	 Sheet flood Sheet wash – pedogenically modified Palustrine flat Braided stream Fluvial under semi-arid, pedogenically modified and flat-palustrine system (Tandon et al., 1990, 1995)	 Overbank Channel Overbank Lacustrine Flood plain Back-swamp (Mohabey, 1996a,b; Mohabey and Samant, 2005) 	 Palustrine environment Fluvial or mostly lacustrine environment (Sarkar et al., 1991; Singh and Tandon, 2004) 	 Gravity flow Flood plain Flood plain Flood plain Pedogenic Lacustrine Fluvial Fluvial Point bar Channel floor Srivastava and Mankar, 2010, 2012b; 2013a) 	



Tab. 4. Bones and eggs of dinosaurs reported from various inland basins of Lameta sedimentation in central and western India.

Bones				Nests/Eggs			
Order	Earlier identified species	New defined species	Oogenus	Earlier identified oospecies	New defined oospecies		
	Balasinor-Jhabua Basin						
Sauropoda	 Antarctosaurus septentrionalis (Mathur and Pant, 1986) Titanosaurus rahioliensis (Mathur and Srivastava, 1987) 	Jainosaurus septentrionalis (Wilson et al., 2003)	Megaloolithus	 M. rahioliensis (=cylindricus), M. phensaniesis, M. khempurensis (=megadermu), M. kachchhensis, M. dhoridungriensis, M. dhoridungriensis, M. megadermus M. balasinorensis (junior synonym of baghensis), Phensaniensis (junior synonym of mohabeyi), Problematica (?) (Mohabey, 1996b; Khosla and Sahni, 1995; Vianey-Liaud et al., 2003) 	 M. cylindricus, M. jabalpurensis M. megadermus Fusioolithus bagbensis (Fernández and Khosla, 2015) 		
Theropods	 Rajasaurus narmadensis (Wilson et al., 2003) Rahiolisaurus gujaratensis 	R <i>ajasaurus narmadensis</i> (Wilson et al., 2003)	Ellipsoolithus				
			Jabalpur Basin				
Sauropods	 Titanosaurus indicus (Lydekker,1877) Antarctosaurus septentrionalis (Huene and Matley, 1993) 	 Titanosaurus colberti (Isisaurus colberti) (Wilson and Upchurch, 2003) Jainosaurus septentrionalis (Hunt et al., 1994) 	Megaloolithus8	 M. dhoridungriensi, M. cylindricus, M. jabalpurensis, M. matleyi (Junior synonym of jabalpurensis) M. phensaniensis (Junior synonym of matleyi) M. mohabeyi, M. baghensis, M. dholiyaensis, M. padiyalensis (Junior synonym of Mohabeyi), M. dhoridungriensis (Mohabey, 1996b; Khosla and Sahni, 1995; Vianey-Liaud et al., 2003) 	 M. cylindricus, M. jabalþurensis M. megadermus F. baghensis (Fernández and Khosla, 2015) 		

Mankar and Srivastava



Tab. 4. (cont.) Bones and eggs of dinosaurs reported from various inland basins of Lameta sedimentation in central and western India.

Bones			Nests/Eggs			
Order	Earlier identified species	New defined species	Oogenus	Earlier identified oospecies	New defined oospecies	
		Jabalpur Basin (co	nt.)			
Theropods	 Indosuchus raptorius (Huene and Matley, 1993) Indosuchus matleyi (Huene and Matley, 1993) Lametasaurus indicus (Matley, 1924) Composuchus solus (Huene and Matley, 1993) Laevisuchus indicus (Huene and Matley, 1993) Lubbalppuria tenuis (Huene and Matley, 1993) Dryptosauroides (?)grandis (Huene and Matley, 1993) Ornithomimoides mobilis (Huene and Matley, 1993) Ornithomimoides barasimlensis (Huene and Matley, 1993) Ornithogoniosaurus Metleyi (Das-Gupta, 1930) Coeluroides largus (Huene and Matley, 1993) Brachypodosaurus gravis (Huene and Matley, 1993) 	 Indosuchus raptorius (Huene and Matley, 1993) Indosuchus matleyi (Huene and Matley, 1993) Laevishuchus indicus (Huene and Matley, 1993) 	Ellipsoolithus	E. khedaensis (Mohabey, 1996b)	E. khedaensis (Mohabey, 1996b)	
		Nand-Dongargaon l	Basin			
Sauropods	 Titanosaurus indicus (Lydekker,1877) Titanosaurus blandfordi (Lydekker,1879) Titanosaurus colberti (Jain and Bandyopadhyay, 1997) Laplatasurus madagascariensis (Huene and Matley, 1933) 	<i>Titanosaurus colberti (Isisaurus colberti)</i> (Wilson and Upchurch, 2003)	Megaloolithus	1. <i>M. matleyi</i> 2. <i>M. megadermus</i> (Mohabey, 1996b)	 M. megadermus M. jabalpurensis (Fernández and Khosla, 2015) 	
Theropods			Ellipsoolithus			
Salbardi-Belkher Basin						
Sauropods		<i>Titanosaurus colberti (Isisaurus colberti)</i> (Wilson and Upchurch, 2003)	Megaloolithus		Megaloolithus (oogenus) (Srivastava and Mankar, 2015b)	

Mankar and Srivastava

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REVIEW PAPER

and fragments of Megaloolithus dhoridungriensis, Megaloolithus cylindricus, Megaloolithus jabalpurensis, Megaloolithus matleyi (junior synonym of jabalpurensis), Megaloolithus phensaniensis (junior synonym of matleyi), Megaloolithus mohabeyi, Megaloolithus baghensis, Megaloolithus dholiyaensis, Megaloolithus padiyalensis (junior synonym of mohabeyi), and M. dhoridungriensis of the Megaloolithidae oofamily, and Elipsoolithus khedaensis of the Elipsoolithidae oofamily (Mohabey, 1983, 1990, 1996b, 1998, 2000, 2001, 2005; Brookfield and Sahini, 1987; Joshi, 1995; Khosla and Sahini, 1995; Vianey-Liaud et al., 2003). Recently, Fernández and Khosla (2015) revised dinosaurian eggs of the Megaloolihidae oofamily reported from the Lameta Formation of India and also reexamined all the nine Megaloolihus oospecies reported by Vianey-Liaud et al. (2003) and confined them only in four oospecies viz., *M. cylindricus, M. jabalpurensis, M. megadermus* and *Fusioolithus baghensis* (Table 4).



Fig. 4. Schematic model of the depositional environment for Lameta sediments: A) deposition of arenaceous sediments as channel floor and point-bar deposits over Gondwana sediments, B) development of non-perennial lakes allowing deposition of fine-grained argillaceous sediments, C) large sheet-like deposits have taken place due to overflooding in the entire area including lakes, and D) volcanic activity causing reduction of sites for fluvial-lacustrine deposition. Pedogenic activities might have taken place at any stage but more pronounced during B stage (revised after Srivastava and Mankar, 2015 a).

6.2. Nand-Dongargaon (N-D) inland basin

Mohabey (1996 a, b) and Mohabey and Samant (2005) have made detailed investigations of lithological set-up and dinosaurian remains of N-D basin. This basin covers an area about of 700 km² constituted together by parts of Nagpur and Chandrapur districts of Maharashtra (Mohabey et al., 1993; Mohabey, 1996 a, b; Mohabey and Samant, 2005). Good successions of Lameta Formation are exposed at Pisdura, Dongargaon, Nand, Kotabala of Chandrapur district and Rajulwari, Pahami and Shivapur of Nagpur district (Mohabey and Samant, 2013; Khosla and Verma, 2015). These successions rest unconformably over the Precambrian basement or, the Gondwana rocks. Towards the northern part of basin i.e., in Nand locality, it rests over the Precambrian granite and schist and is marked by 4 m thick lithocolumn of sandstone, which is conglomeratic at the base and calcretised in upper part (Mohabey, 1996 a). In Dongargaon-Pisdura, towards the southern extreme, the Lameta succession comprises of red and green, nonlaminated silty clay that unconformably rests over the Precambrian granite and schist or, the Kamthi Formation (Mohabey et al., 1993). It is argillaceous in nature forming a column of about 6 m thickness having interbeddings of sandstone (Fig. 2). Overlying sediments are yellow and cream-colored laminated clays and shales with thin beds of limestones and marlites. The top 4 m thick column, of which, upper part shows calcretization (Mohabey and Samant, 2005, 2013; Khosla and Verma, 2015). Mohabey (1996 a) based on detailed lithological characteristics, identified four major lithounits viz., i) red and green silty clays associated with sandstones, ii) channel related sandstone (trough and cross-bedded), which are calcretised in the upper part, iii) yellow laminated clays and shales interbedded with limestone, marlite and septarian concretion bands and, iv) gray nodular marls. Based on lithological and sedimentological attributes and faunal content, the above mentioned lithounits are interpreted to be deposited under



i) channel, ii) overbank, iii) paludal and iv) limnic environments respectively (Mohabey et al., 1993; Mohabey and Udhoji, 1993; Mohabey, 1996 a; Singh and Tandon, 2004). The lacustrine facies is developed in the southern sector of N-D basin which have been further studied in detail revealing sedimentation pattern, climatic control and sediments cyclicity (Mohabey and Samant, 2005).

6.2.1. Dinosaur remains

Sauropod remains belonging to *T. indicus, T. blandfordi, T. colberti, Isisaurus colberti* and *Laplatasurus madagascariensis* were reported from this basin (Lydekker, 1890; Matley, 1921; Chakravarti, 1934, 1935; Berman and Jain, 1982; Jain and Bandhopadhya, 1987; Vianey-Liaud et al., 1987; Jain and Bandhopadhyay, 1997; Carrano et al., 2010). However, according to Wilson and Upchurch (2003) only *T. colberti* (= *Isisaurus colberti*) would be a valid genus (Table 4).

6.2.2. Nests and Eggs

Megaloolithus matleyi and M. megadermus oospecies of the Megaloolithus oofamily were reported from this basin (Jain and Sahni, 1985; Khosla and Sahni, 1995; Mohabey, 1998, 2000, 2001; Vianey-Liaud et al., 2003); however, according to Fernández and Khosla (2015), only M. megadermus oospecies would be a valid species. Chelonian and spherolithidae eggs have also been reported (Mohabey, 1996b, 1998) (Table 4).

6.3. Balasinor-Jhabua inland basin

The Lameta Formation, outcropping at Balasinor and Rahioli areas of Kheda and Panchmahal districts of Gujarat were heavily explored for dinosaur bones and eggs. At Balasinor, the 4 to 12 m thick succession is unconformably developed over the Godhra Granites/Phyllites of Aravalli Supergroup (Fig. 2). The lower 2 to 6 m thick succession is conglomeratic in nature and shows channel and scour-fill structures with pebbles, quartz veins, cherty quartzite bounded together by calcareous and siliceous cements. This horizon also shows the preservation of abundant dinosaur bones. The bedding characteristic is not distinct; however, locally, the size gradation of the clasts has been reported. This level grades to 2 m thick, pebbly, poorly sorted sandstone. The sandstone gradually becomes fine grained in upward direction and later grades to limestone which is gray to brown in color and siliceous in nature. The limestone is highly variable in nature i.e., mottled, nodular and brecciated with lithoclasts, peloids and subangular to subrounded detrital quartz. Mohabey (2001) considered it a deposit of palustrine environment which is based on its fine grained, massive and micritic nature, in addition, to the presence of discontinuous cracks and bioturbation. At Rahioli, the sandy calcrete is comparatively well developed and entombs almost complete eggs and nests of dinosaurs (Mohabey, 1984; Srivastava et al., 1986; Sahni et al., 1994; Loyal et al., 1996, 1999; Fernández and Khosla, 2015; Khosla and Verma, 2015). The depositional environments range mostly from sheet wash to palustrine. The oxygen and carbon isotopic analysis are also suggestive of fluvial or mostly lacustrine environment of deposition, particularly for the egg-bearing horizons (Sarkar et al., 1991; Singh and Tandon, 2004).

6.3.1. Dinosaur remains

Bones of the Antarctosaurs septentrionalis and the Rajasaurus narmadensis belonging to the sauropod and theropod respectively (Jain and Bandhopadhya, 1997; Wilson and Upchurh, 2003; Wilson et al., 2003) (Table 4).

6.3.2. Nests and Eggs

Megaololithus rahioliensis (=cylindricus), Megaololithus phensaniesis, Megaololithus khempurensis (=megadermus), Megaololithus kachchhensis, M. dhoridungriensis, M. megadermus and M. balasinorensis (junior synonym of baghensis), M. phensaniensis (junior synonym of mohabeyi), Problematica (?) reported earlier from this basin (Dwivedi et al., 1982; Mohabey, 1983, 1991; Khosla and Sahni, 1995; Loyal et al., 1999; Vianey-Liaud et al., 2003) are now redefined as Megaololithus cylindricus, M. jabalpurensis, M. megadermus and Fusioolithus baghensis by Fernández and Khosla, 2015 (Table 4).

6.4. Sagar (Saugor) and Amarkantak-Ambikapur inland basins

Both basins are comparatively less studied and devoid of adequate information in the original paper of Mohabey (1996 a). The Sagar basin, laying in the NW of Jabalpur basin, was categorized on the basis of a north-west elongated patches of Lameta sediments in its geographical areas. This succession consists of gritty sandstone, calcareous clay, chert and mottled limestone having pebbles of jasper and chert. The Amarkantak-Ambikapur basin lying in the uplift part constitutes comparatively very large geographical area; however, lacks sufficient details (Mohabey, 1996 a).

7. Discussion

Mohabey (1996 a) proposed five inland basins viz., i) Nand-Dongargaon, ii) Jabalpur, iii) Sagar, iv) Ambikapur-Amarkantak and v) Balasinor-Jhabua, in which, the deposition of the Lameta sediments had taken place. The fluvial-lacustrine sediments of Lameta, exposed in these inland basins, have a remarkable similarity in their architectural set-up (Fig. 2). The authors of the present work have recently explored and studied a few isolated patches of Lameta sediments exposed at Bairam, Belkher, Pandhari and Salbardi areas along with 3-4 minor exposures, covering together a large geographical area in wide spread Deccan



REVIEW PAPER

Trap basaltic country. The successions in these areas are fluvio-lacustrine in nature and preserve evidences of dinosaur, including bones, coprolites, nests and eggs. Based on the comparative study of the lithological set-ups and dinosaurian remains of this basin with those of other five inland basins reported earlier, the present authors have already proposed it as a new inland basin for Lameta sedimentation and named Salbardi-Belkher inland basin (Fig. 1). In this basin, the commencement of Lameta sedimentation, after the termination of Gondwana sedimentation, is well marked, however, the sediment supply during deposition of both the lithounits have a common source. It is a significant outcome as at other places in Central and Western India, the Lameta sediments lie unconformably over the older basement with pronounced unconformity.

On the basis of facies analysis, it is interpreted that the sedimentation in the new basin took place in a big riverine plain including lakes which was periodically recharged by the river causing shallowing and deepening of the depositional site in entire course of Lameta sediments deposition. This phenomenon occurred repeatedly during the sedimentation, as the column in the lower part is dominantly represented by bedded units of arenaceous and argillaceous facies depicting the change in energy condition of the depositional environment and also the periodic influx of the sediments having high variation in grain size. During the later stage, the basin was mostly shallow and alkaline in nature represented by dominance of calcareous sediments. In the initial phase, the basin was probably shallow in the eastern part and deep in the west, which is evident by the difference in the nature of sediments in both the geographical extremities. The succession, west of Salbardi area, is noticed to be of arenaceous nature in the lower part that also contains calcrete horizons revealing the shallow nature of the depositional site as compared to the dominance of argillaceous sediments at Belkher and further Pandhari in the extreme west showing comparatively deeper condition.

Comparing it with the other five inland basins as proposed by Mohabey (1996 a), Mankar and Srivastava (2015) infer that the Salbardi-Belkher inland basin was almost similar in size as of Jabalpur; however, elongated in east-west directions. Like others, in this basin also, the deposition of the sediments has a basinal control; however, the thickness of the column is comparatively larger. The dinosaur bones and eggs-bearing horizons are also correlated with similar horizons of various localities of other basins viz., Raholi, Jabalpur and Nand-Dongargaon, that also represent similar deposition environment set-up i.e., channel to point bar subenvironment under sub and climatic condition.

Taking the newly identified Salbardi-Belkher basin (Mankar and Srivastava, 2015) into consideration, the palaeogeographic set-up during Lameta sedimentation is now been revised with extended geographical limit. The climatic condition in the newly identified basin area was also favorable for survival of dinosaurs and laying their eggs as of Jabalpur, N-D and Balasinor-Jhabua basins hence, the same condition is now extrapolated in the west of adjacent N-D basin.

8. Conclusion

In view of the studies carried out for lithological set-up and dinosaurian remains of newly identified Salbardi-Belkher fluvial inland basin, and its comparison with earlier established five co-eval depositional basins, the following conclusions have been drawn for the Lameta Formation:

- The fluvio-lacustrine Lameta sediments exposed in central and western India, earlier considered to be deposited in only five inland basins viz., i) Nand-Dongargaon, ii) Jabalpur, iii) Sagar (Saugor), iv) Ambikapur-Amarkantak and v) Balasinor-Jhabua are now added with an additional basin viz., Salbardi-Belkher making a total of six basins of deposition.
- 2. The newly identified Salbardi-Belkher inland basin lies in western side of Nand-Dongargaon basin and preserves well developed fluvio-lacustrine Lameta successions exposed at Bairam, Belkher, Pandhari and Salbardi areas besides 3-4 small exposures. Skeletal remains, coprolites, eggs, eggs-nests of dinosaurs are also preserved.
- 3. A comparative study of the lithological set-ups of earlier known five inland basins of Lameta sediments with those of Salbardi-Belkher basin exhibits similar depositional set-up; however, in the newly established basin, an additional lithounit of gravity flow represented by intraformational brecciated limestone has also been identified at the top.
- 4. Similarly, the overall comparison of dinosaurian remains show that the *Titanosaurus colberti* (= *Isisaurus colberti*), species of sauropod, reported form N-D and Balasinor-Jhabua basin, also existed in Salbard-Belkher basin. The eggs and nests of *Megaloolithus* reported from the Salbardi-Belkher basin is also comparable with Balasinor-Jhabua, Jabalpur and N-D basins at oogenus level.
- 5. Though, the Salbardi-Belkher basins is very close to N-D basin having almost similar fluvio-lacustrine depositional set-up but, differs remarkably in their lithological set-ups and sub-environments of deposition showing different basinal set-ups in both basins.
- 6. The inclusion of new inland basin with separate geographical identity reflects a revised paleogeographic set-up having a westward extended limit of Lameta sedimentation and dinosaur inhabitation as compared to the paleogeographic limits proposed earlier on the basis of only five inland basins.



REVIEW PAPER

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